

Provenance of late Paleozoic sedimentary rocks in eastern Kazakhstan: Implications for the collision of the Siberian margin with the Kazakhstan collage

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ARTICLE INFO

Keywords:

Central Asian Orogenic Belt
Char Shear Zone
Zharma-Saur Arc
Irtysch-Zaisan Complex
Detrital zircon

ABSTRACT

To unravel when, where, and how multiple arc systems were amalgamated is crucial for understanding the accretion processes of the Central Asian Orogenic Belt. Here we focus on the collision of the Siberian margin and the Kazakhstan collage along the Char Shear/Suture Zone in eastern Kazakhstan. We conducted U–Pb dating and Hf isotope analysis for detrital zircons of clastic rocks across the collisional zone of the Siberian margin (the Irtysch-Zaisan Complex) and the Kazakhstan collage (the Zharma-Saur Arc and the Char Zone). Our results show that the Irtysch-Zaisan Complex carries abundant early Paleozoic detrital zircons and minor Precambrian zircons, but there is a lack of such zircon populations in the Zharma-Saur Arc and the Char Zone. Devonian to Carboniferous detrital zircons appear in all sampling tectonic units, in which $\epsilon_{\text{Hf}}(t)$ values of Carboniferous zircons from the Zharma-Saur Arc and the Char Zone are characterized by a narrow range (~10–17) that is different from a relatively wider range of $\epsilon_{\text{Hf}}(t)$ values of ~5–18 for the Carboniferous zircons in the Irtysch-Zaisan Complex. The distinct sedimentary provenances of the Irtysch-Zaisan Complex and the Zharma-Saur Arc/Char Zone confirm that their tectonic boundary lies along the Char Shear Zone in eastern Kazakhstan. As the Late Carboniferous sedimentary rocks from the Zharma-Saur Arc and the Char Zone do not receive detritus from the Siberian margin, we consider that the Ob-Zaisan Ocean between the Siberian margin and the Kazakhstan collage was closed later than the deposition of these sedimentary rocks (<321 Ma).

1. Introduction

The Central Asian Orogenic Belt (CAOB), as a typical accretionary orogen, occupies most of the central Asian interior. The accretion processes and mechanisms of this giant orogenic system has attracted widespread attention in the past decades (Zonenshain et al., 1990; Şengör et al., 1993; Windley et al., 2007; Wilhem et al., 2012; Li et al., 2013; Han et al., 2015; Xiao et al., 2015, 2018; Li et al., 2019). An archipelago paleogeographic configuration has been proposed for the CAOB during the Paleozoic, in which multiple Andean-, Japanese-, Alaska-, Mariana-type arc systems were developed synchronously (Xiao et al.,

2010, 2013; Safonova et al., 2011, 2018; Wilhem et al., 2012). These arc systems were amalgamated following the closure of the Paleo-Asian Ocean (PAO) in the latest Paleozoic, which was accompanied by the occurrence of several large-scale of strike-slip shear zones (Fig. 1a; Qu and Zhang, 1991, 1994; Şengör et al., 1993; Allen et al., 1995; Shu et al., 1999; Travin et al., 2001; Laurent-Charvet et al., 2003; Buslov et al., 2004; Li et al., 2017; Han and Zhao, 2018; Hu et al., 2020). This collisional phase significantly obliterated earlier accretion-related records, and modified the architecture of the CAOB. Therefore, it is crucial firstly to unravel when, where and how multiple arc systems were amalgamated in order to understand the accretion mechanisms of the CAOB.

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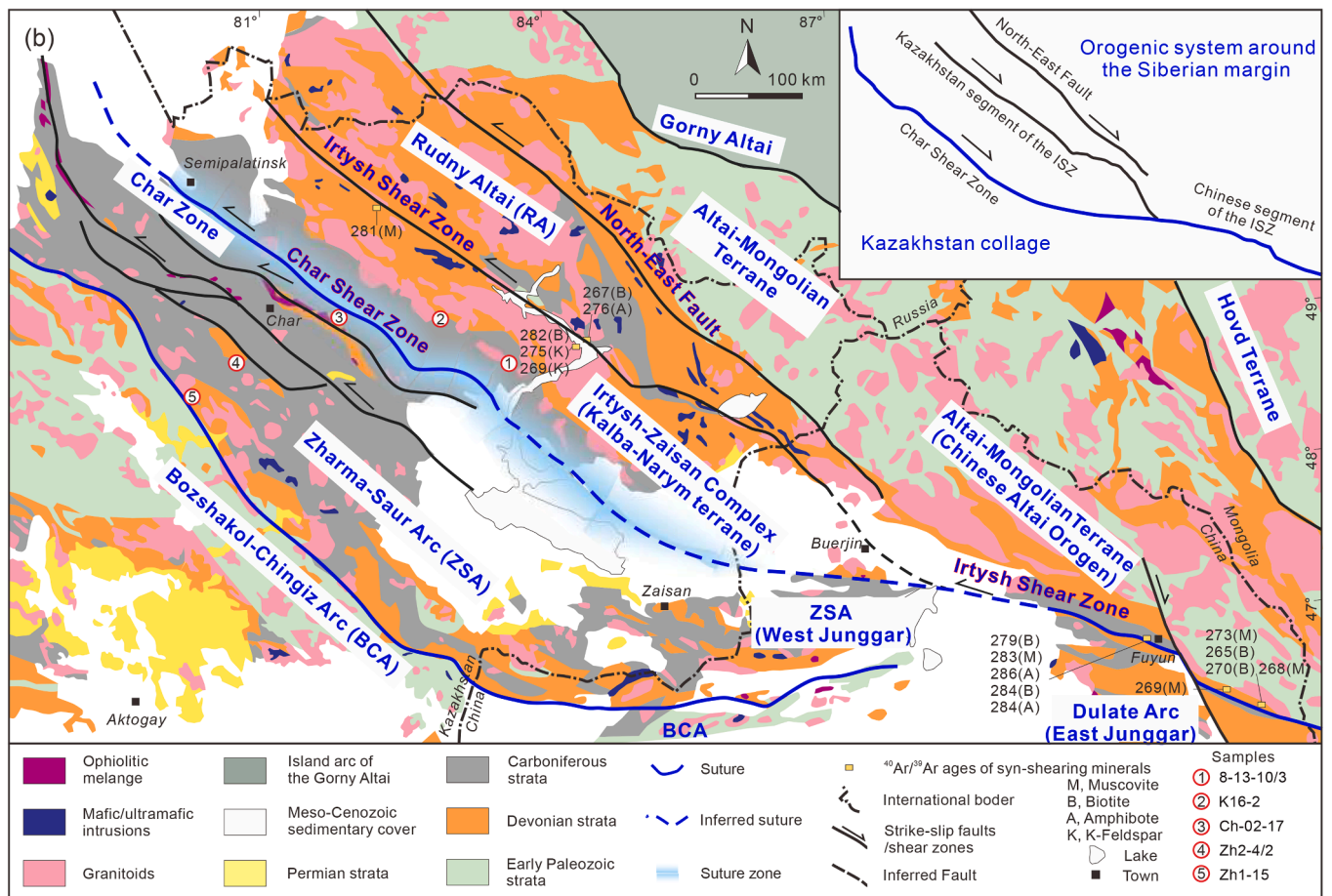
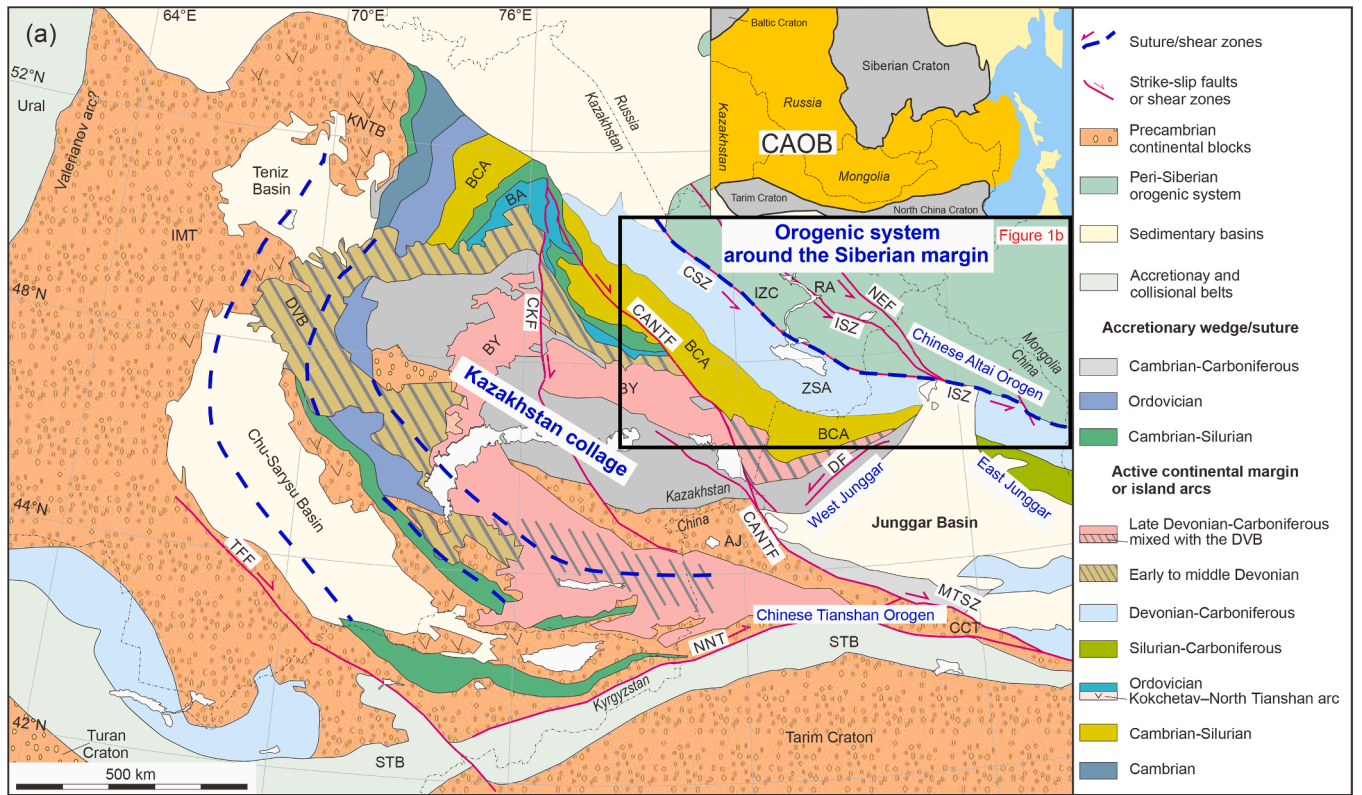
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<https://doi.org/10.1016/j.jseaes.2021.104978>

Received 30 May 2021; Received in revised form 10 October 2021; Accepted 14 October 2021

Available online 25 October 2021

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Fig. 1. (a) The tectonic map of the western CAO (modified after Windley et al., 2007; Li et al., 2018; and Hu et al., 2020). The inset image shows the location of the CAO (after Jahn, 2004). (b) A simplified geological map along the Irtysh-Char Shear Zone that marks the tectonic boundary of Siberian margin and the Kazakhstan collage (after Li et al., 2008). The inset figure highlights the major shear zones/faults along the collisional front of the Siberian margin and the Kazakhstan collage. Note that the Char Shear Zone that represents the tectonic boundary of the Siberian margin and the Kazakhstan collage in eastern Kazakhstan, merges into the Irtysh Shear Zone in NW China, where the Irtysh Shear Zone marks the boundary of the Siberian margin and the Kazakhstan collage. The $^{40}\text{Ar}/^{39}\text{Ar}$ ages of syn-shearing minerals along sinistral strike-slip shear zones are from Travin et al., 2001, Laurent-Charvet et al., 2003, Li et al., 2017, and Hu et al., 2020. Abbreviations, RA: Rudny Altai; IZC: Irtysh-Zaisan Complex; ZSA: Zharma-Saur Arc; BCA: Bozshakol-Chingiz Arc; BYA: Balkhash–Yili Arc; DVB: Devonian Volcanic Belt; NEF: North-East Fault; ISZ: Irtysh Shear Zone; CANTF: Chingiz-Alakol-North Tianshan Fault; MTSZ: Main Tianshan Shear Zone.

The NW-SE Char Shear Zone (also termed Gornostaev Shear Zone by Şengör et al., 1993) in eastern Kazakhstan has been interpreted to be a major suture zone within the CAO, which represents the tectonic boundary between the orogenic system around the Siberian margin and the Kazakhstan collage (Fig. 1a; Buslov et al., 2001, 2004; Volkova et al., 2008; Safonova et al., 2012, 2018; Kuibida et al., 2016). It extends southeastward into NW China, where it is merged with the Irtysh Shear Zone (Fig. 1b; Laurent-Charvet et al., 2003; Li et al., 2017). Structurally, both Char and Irtysh shear zones show sinistral strike-slip kinematics (Buslov et al., 2003; Vladimirov et al., 2008; Li et al., 2015; Hu et al., 2020), which have been considered to result from the oblique convergence of the orogenic system around the Siberian margin and the Kazakhstan collage after the closure of the Ob-Zaisan Ocean (a branch of the Paleo-Asian Ocean that is also referred to as the Irtysh Ocean) (Qu and Zhang, 1991, 1994; Laurent-Charvet et al., 2003; Li et al., 2015, 2017; Hu et al., 2020). Available studies show that both the ocean plate stratigraphy (OPS) and arc-related fragments occur along the Char Shear Zone (Safonova et al., 2012, 2018, 2021). However, the initial collisional time of the two orogenic systems has been controversial. In eastern Kazakhstan, Glorie et al. (2012) proposed that the Ob-Zaisan Ocean was closed in the Early Carboniferous, on a basis of the interpreted syn-collisional origin of ~338 Ma intrusion. The subsequent geochemical studies for ~323–317 Ma igneous rocks by Kuibida et al. (2016) and Khromykh et al. (2019) also favor a syn-collisional origin for these rocks. In contrast, Zhang et al. (2020) proposed a subduction-related origin for 330–318 Ma igneous rocks to the north of the Char Shear Zone, indicating that the northward subduction of the Ob-Zaisan oceanic plate probably lasted until ~318 Ma. Such inconsistency shows the limitation of geochemical affinities of igneous rocks in tracing the tectonic transition associated with the closure of the oceanic basin. Farther southeast in NW China, detrital zircon data show that the Ob-Zaisan Ocean existed until the Late Carboniferous given that there is no exchange of detrital zircons for the orogenic system around the Siberian margin (Chinese Altai) and the Kazakhstan collage (West Junggar) in this period (Choulet et al., 2012a; Li et al., 2017; Song et al., 2020a). It remains enigmatic whether the Ob-Zaisan Ocean in NW China and eastern Kazakhstan share a similar tectonic history (e.g. synchronous/diachronous closure along the strike).

In this paper, we present new detrital zircon U–Pb and Hf isotope data from the Irtysh-Zaisan Complex, the Char Zone and the Zharma-Saur Arc in eastern Kazakhstan, with an aim to constrain the collisional time of the orogenic system around the Siberian margin and the Kazakhstan collage along the Char Shear Zone. Our results show that the Ob-Zaisan Ocean between the Siberian margin and the Kazakhstan collage in eastern Kazakhstan existed until ~321 Ma, similarly as the evolution of eastern segment of this oceanic basin in NW China.

2. Geological setting

The Char Shear Zone in eastern Kazakhstan marks the tectonic boundary of the orogenic system around the Siberian margin and the Kazakhstan collage (Fig. 1b; Buslov et al., 2001; Vladimirov et al., 2008; Volkova et al., 2008; Safonova et al., 2012; Kuibida et al., 2016; Safonova et al., 2018). The former was built by the accretion of a series of microcontinental blocks, island arcs, and accretionary complexes along the Siberian margin from Neoproterozoic to late Paleozoic (Wilhem et al., 2012; Xiao et al., 2015; Li et al., 2019), while the latter was

predominantly developed via the amalgamation of microcontinental blocks (Kazakhstan microcontinent) and arc terranes in the early Paleozoic, followed by late Paleozoic oroclinal bending (Fig. 1a; Khain et al., 2003; Windley et al., 2007; Abrajevitch et al., 2008; Xiao et al., 2010; Degtyarev, 2011; Choulet et al., 2012b; Levashova et al., 2012; Rolland et al., 2013; Li et al., 2018). A ~100 km long ophiolitic belt (hosting OPS units) occurs within the collisional zone of the Siberian margin and the Kazakhstan collage (Char Zone; Fig. 1b), with high-pressure rocks (e.g. blueschist and eclogite that are metamorphosed from oceanic basalts) incorporated into the ophiolitic mélange (Buslov et al., 2001; Vladimirov et al., 2008; Volkova et al., 2008; Safonova et al., 2012, 2018).

2.1. Tectonic elements around the Siberian margin in eastern Kazakhstan

In eastern Kazakhstan, three ~NW-SE tectonic units of the Altai-Mongolian Terrane, the Rudny Altai, and the Irtysh-Zaisan Complex were developed along the Siberian margin (Fig. 1b). The northernmost unit of the Altai-Mongolian Terrane extends from eastern Kazakhstan, through NW China (Chinese Altai), to western Mongolia (Fig. 1b), and it mainly comprises early Paleozoic meta-sedimentary/volcanic rocks (Buslov et al., 2001; Badarch et al., 2002; Windley et al., 2002; Cai et al., 2011a). This terrane has been interpreted to be an accretionary complex (Long et al., 2008, 2012; Jiang et al., 2012; Chen et al., 2014, 2015), which together with the coeval magmatic arc (Ikh-Mongol arc) in western Mongolia constitutes an Andean-type subduction system along the Siberian margin (Bold et al., 2016; Soejono et al., 2017; Janousek et al., 2018; Li et al., 2019). The Altai-Mongolian Terrane is bounded with the Rudny Altai farther south by a sinistral strike-slip fault of the North-East Fault (Fig. 1b). The Rudny Altai has been interpreted to be a back-/intra-arc extensional basin that was developed along the southern margin of the Altai-Mongolian Terrane (Buslov et al., 2004; Lobanov et al., 2014; Kuibida et al., 2020). The oldest rocks in the Rudny Altai are Silurian to Devonian meta-sedimentary rocks (metasandstones, schists, and phyllites), which are overlain by Devonian to Carboniferous rhyolite-basalt lava, volcanic breccia, tuff, conglomerate, sandstone and minor limestone (Buslov et al., 2001; Buslov et al., 2004; Kruk et al., 2011; Lobanov et al., 2014). Farther south, the Irtysh-Zaisan Complex (also termed the Kalba-Naryn terrane, Buslov et al., 2001) is separated from the Rudny Altai by the Irtysh Shear Zone (Şengör et al., 1993; Buslov et al., 2004; Glorie et al., 2012). The Irtysh-Zaisan Complex contains migmatite, amphibolite, gneiss, schist, metasandstone, quartzite, marble, as well as a small amount of metamorphosed basalt and chert (Şengör et al., 1993; Buslov et al., 2004; Briggs et al., 2007; Chen et al., 2019). Its counterpart in NW China, has been interpreted as a Devonian to Carboniferous accretionary complex consisting of oceanic plate stratigraphy and continent-marginal sedimentary rocks (O'Hara et al., 1997; Briggs et al., 2007; Li et al., 2015, 2017; Xiao et al., 2015; Chen et al., 2019).

The Altai-Mongolian Terrane, the Rudny Altai, and the Irtysh-Zaisan Complex in eastern Kazakhstan are voluminously intruded by late Paleozoic granitoids (Fig. 1b; Glorie et al., 2012; Khromykh et al., 2016; Kuibida et al., 2019, 2020; Zhang et al., 2020; Kotler et al., 2021). The Devonian to Carboniferous magmatism was considered to be associated with the northward subduction of the Ob-Zaisan oceanic plate (Buslov et al., 2004; Safonova, 2014; Kuibida et al., 2020; Zhang et al., 2020), while the Permian intrusions were likely related to the collision of the

Siberian margin and the Kazakhstan collage (Vladimirov et al., 2008; Kuibida et al., 2009, 2019).

2.2. Island arcs of the Kazakhstan collage in eastern Kazakhstan

The Kazakhstan collage shows a U-shaped geometry in a map view (Kazakhstan Orocline; Şengör et al., 1993), which is delineated by two curved magmatic arc systems of the Devonian Volcanic Belt and the Late Devonian to Carboniferous Balkhash–Yili Arc (Fig. 1a; Levashova et al., 2003, 2012; Windley et al., 2007; Abrajevitch et al., 2008; Degtyarev, 2011; Bazhenov et al., 2012; Li et al., 2018). Along the northern limb of the orocline, both arc systems were developed over an early Paleozoic island arc of the Bozshakol-Chingiz Arc (Fig. 1a) that mainly contains Cambrian to Silurian mafic to felsic volcanic rocks, pyroclastic rocks, and chert (Feng et al., 1989; Degtyarev and Ryazantsev, 2007; Chen et al., 2010; Shen et al., 2015), and Late Silurian to Carboniferous volcanic rocks and carbonate (Degtyarev and Ryazantsev, 2007; Wei et al., 2009). The late Paleozoic intra-oceanic island arc of the Zharma-Saur Arc occurs to the north of the Bozshakol-Chingiz Arc (Fig. 1a; Safonova et al., 2017). It is dominated by Devonian to Carboniferous volcanic rocks, limestone and clastic rocks (Zhu and Xu, 2006; Vladimirov et al., 2008; Chen et al., 2010), which are intruded by late Paleozoic mafic to felsic plutons (Han et al., 2006; Vladimirov et al., 2008; Zhou et al., 2008; Chen et al., 2010; Zhang et al., 2015a; Liu et al., 2018). Farther north, the Char Zone is dominated by mafic to andesitic volcanic

rocks, tuff and sandstone, which is mixed with the ophiolitic mélangé and is affected by a series of sinistral strike-slip faults (Char Shear Zone; Fig. 1b). The volcanic rocks within the Char Zone show Late Devonian to Early Carboniferous zircon U–Pb ages and positive $\epsilon_{\text{Nd}}(t)$ values (3–9) (Safonova et al., 2018). The greywackes of the Char Zone are characterized by Carboniferous detrital zircon age peaks, positive zircon $\epsilon_{\text{Hf}}(t)$ values (mostly in the range of 5 to 14) and whole-rock $\epsilon_{\text{Nd}}(t)$ values (5–8) (Safonova et al., 2021).

3. Zircon U–Pb geochronology and Hf isotope analysis

3.1. Sample description

Five samples were collected for detrital zircon U–Pb dating and Hf isotope analysis with an aim to constrain the provenance variation across the Char Shear Zone in eastern Kazakhstan. All sample locations are illustrated in Fig. 1b. Two samples (8–13–10/3 and K16–2; Fig. 1b) are from the Irtysh-Zaisan Complex to the north of the Char Shear Zone. Sample 8–13–10/3 (GPS coordinate: 49°06′11.11″ N, 83°20′22.15″ E) is a medium-grained sandstone with oriented lens-shaped quartz and feldspar to define a foliation (Fig. 2a). The other sample from the Irtysh-Zaisan Complex (K16–2; GPS coordinate: 49°30′47.40″ N, 82°48′15.00″ E) is a mica quartz schist, with a mineral assemblage of quartz, muscovite and minor biotite (Fig. 2b).

Two samples (Zh2–4/2 and Zh1–15; Fig. 1b) were collected from

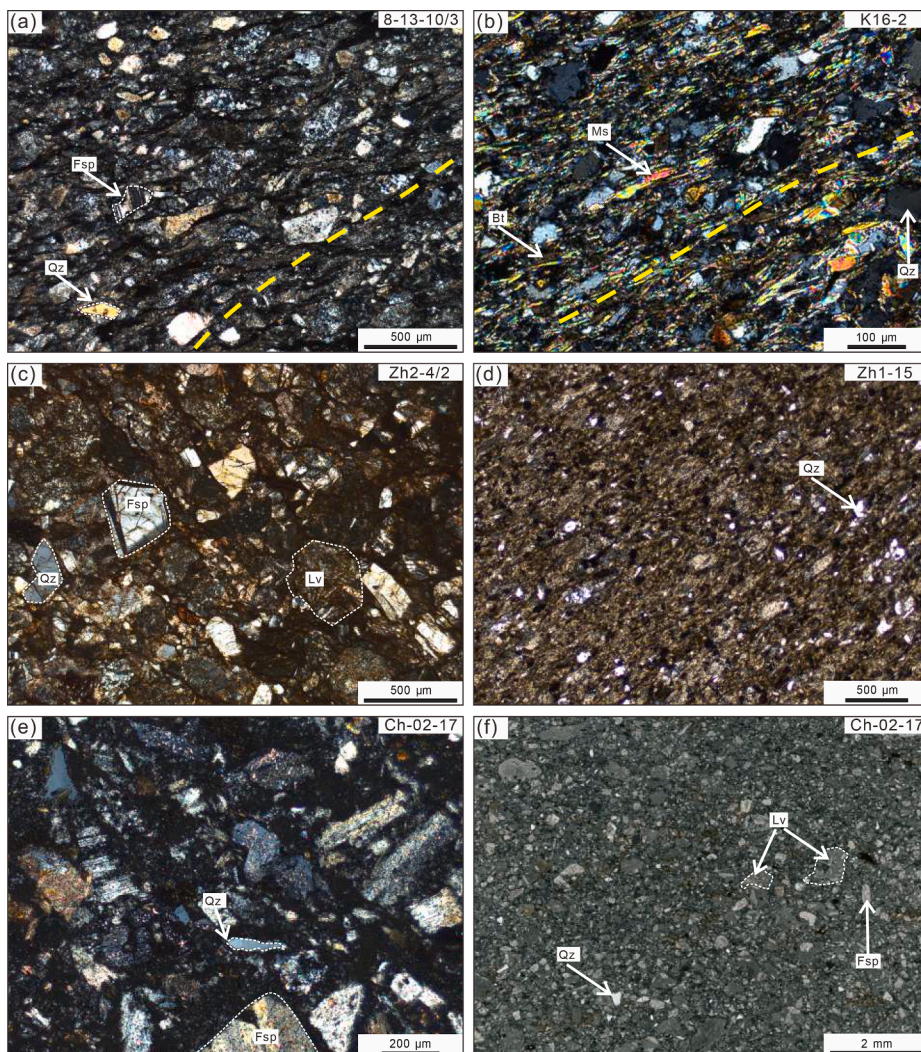


Fig. 2. Photographs of samples for zircon U–Pb Hf isotope analysis. (a) Medium-grained sandstone from the Irtysh-Zaisan Complex, with a secondary foliation indicated by yellow dashed line. (b) Mica quartz schist from the Irtysh-Zaisan Complex, with preferred alignment of muscovite and biotite to define a foliation (yellow dashed line). (c) Medium-grained sandstone from the Zharma-Saur Arc, which contains quartz, feldspar, and minor calcite and volcanic lithic fragments. (d) Siltstone from the Zharma-Saur Arc, which is mainly composed of quartz, feldspar, and opaque minerals. (e–f) Fine-grained sandstone, which was collected along the Char Zone, containing feldspar, volcanic lithic fragments, and minor quartz and opaque minerals. Abbreviations, Qz: quartz; Fsp: feldspar; Ms: muscovite; Bt: biotite; Lv: volcanic lithic fragment. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Carboniferous strata of the Zharma-Saur Arc. Sample Zh2-4/2 (GPS coordinate: 49°13'46.09" N, 80°45'40.08" E) is a medium-grained sandstone that contains quartz, feldspar, and minor calcite and volcanic lithic fragments (Fig. 2c), while sample Zh1-15 (GPS coordinate: 48°56'54.07" N, 80°14'15.03" E) is a moderately sorted siltstone that is composed of quartz, feldspar, and opaque minerals (Fig. 2d).

One additional sample (Ch-02-17; GPS coordinate: 49°35'07.00" N, 81°52'44.04" E) was taken along the northeastern flank of the Char Zone (Fig. 1b). It is a fine-grained sandstone that is composed of sub-angular feldspar, volcanic lithic fragments, quartz, and opaque minerals (Fig. 2e, f).

3.2. Methods

Zircon grains were extracted using traditional heavy liquid and magnetic techniques, and were mounted in an epoxy resin. Cathodoluminescence (CL) images were taken in order to unravel the internal structure of zircons. Zircon U–Pb dating was conducted by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) (Resonetics RESOLUTION S-155 laser + Agilent 7900) at the Guangzhou Institute of Geochemistry, Chinese Academy of Sciences (GIG-CAS). For detailed operating conditions and analytical procedures, see Li et al. (2012). The laser diameter was set as 29 μm for samples 8–13-10/3, K16-2, and Zh2-4/2, and 19 μm for samples Zh1-15 and Ch-02-17. The frequency of the laser was 6 Hz. Zircon 91500 was used as external standards. The content of trace elements in zircon was quantitatively



Fig. 3. Representative CL images of analyzed zircon grains, in which U–Pb ages and $\epsilon_{\text{Hf}}(t)$ values are illustrated. Analyzed spots for U–Pb dating and Hf isotope analysis are indicated by yellow and blue circles, respectively. The black line below each zircon image is for scale (50 μm). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

analyzed with NIST610 as an external calibration reference and Si as an internal standard. The first 20 s was set to detect the gas blank, followed by 50 s laser ablation for each spot signal. The ICPMSDataCal 10.8 and the Isoplot 4.15 were used for performing off-line analyzed signals and calculating ages, respectively (Ludwig, 2003; Liu et al., 2008; Lin et al., 2016).

Hf-in-zircon isotopic analyses of three samples (8–13–10/3, K16–2, and Zh2–4/2) were conducted at the GIG-CAS via using a Neptune Plus MC-ICP-MS (Thermo Scientific), coupled with a RESolution M–50 193 nm laser ablation system (Resonetics). Detailed operating conditions and analytical methods are described in Zhang et al. (2014). In order to improve the instrumental sensitivity, an X skimmer cone was designed in the interface. The single spot size and frequency of the laser were 45 μm and 6 Hz. All isotope signals were detected with Faraday cups under static mode. Each analysis for zircons includes 20 s background signal and 30 s data signal. Isobaric interference of ^{176}Yb and ^{176}Lu on ^{176}Hf was corrected by ^{173}Yb and ^{175}Lu . The corrected natural ratio values of $^{176}\text{Yb}/^{173}\text{Yb}$ and $^{176}\text{Lu}/^{175}\text{Lu}$, are 0.79381 (Segal et al., 2003) and 0.02656 (Wu et al., 2006), respectively. The mass bias factor of Lu is assumed to be the same as that of Yb, which is calculated from the measured $^{173}\text{Yb}/^{171}\text{Yb}$ and the natural ratio of 1.13268 (Chu et al., 2002). The mass bias of $^{176}\text{Hf}/^{177}\text{Hf}$ was normalized to $^{179}\text{Hf}/^{177}\text{Hf} = 0.7325$ with an exponential law. The detailed data reduction procedure is documented in Zhang et al. (2015b). Nineteen Plešovice zircon yielded a weighted mean of $^{176}\text{Hf}/^{177}\text{Hf} = 0.282486 \pm 0.000008$ (2SD), which is consistent within errors with reported in Sláma et al. (2008).

3.3. Results

Representative CL images of analyzed zircons are shown in Fig. 3. The U–Pb and Hf isotope data of zircons are presented in Supplementary Tables 1 and 2. $^{207}\text{Pb}/^{206}\text{Pb}$ ages are used for zircons older than 1000 Ma, and $^{206}\text{Pb}/^{238}\text{U}$ ages are used for zircons younger than 1000 Ma. The values of $\varepsilon_{\text{Hf}}(t)$ were calculated using $^{176}\text{Lu}/^{177}\text{Hf} = 0.0332$ and $^{176}\text{Hf}/^{177}\text{Hf} = 0.282772$ (Blichert-Toft and Albarede, 1997). The depleted mantle line, exhibited in Fig. 5 is based on the present $^{176}\text{Hf}/^{177}\text{Hf} = 0.28325$ and $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ (Griffin et al., 2000).

Detrital zircons from samples 8–13–10/3 and K16–2 within the Irtysh-Zaisan Complex, are predominantly prismatic with variable luminescence in CL images (Fig. 3a, b). Most zircons are 30–200 μm long, and are characterized by oscillatory zoning, which together with Th/U ratio range from 0.22 to 2.96, implying an igneous origin (Supplementary Table 1). One additional zircon grain has a low Th/U ratio (0.06), but it is characterized by oscillatory zoning, indicating a magmatic origin (e.g. Gebauer, 1996; Wu and Zheng, 2004). Total 65 zircon grains were analyzed for sample 8–13–10/3. 61 zircons yield concordant early Paleozoic ages from 313 to 536 Ma, with a major age peak (~ 330 Ma) and three minor age peaks (~ 348 Ma, ~ 376 Ma, and ~ 473 Ma; Fig. 4a, b). The Hf isotope analysis for this group of zircons gives $\varepsilon_{\text{Hf}}(t)$ values ranging from -1.6 to 19.3 (Fig. 5a). Additional 3 analyses give scattering Precambrian ages of 554 Ma, 675 Ma, and 835 Ma. As for sample K16–2, 56 zircon grains yield Paleozoic U–Pb ages (317–533 Ma) with a remarkable age peak at ~ 328 Ma and two minor age peaks of ~ 452 Ma and ~ 488 Ma. These Paleozoic zircons have predominant $\varepsilon_{\text{Hf}}(t)$ values from -6.3 to 20.1 except one 325 Ma zircon grain with -14.8 $\varepsilon_{\text{Hf}}(t)$ value. Additional 4 zircons give Precambrian ages of 619 Ma, 786 Ma, 942 Ma, and 2028 Ma (Fig. 4c, d).

Detrital zircons from samples Zh2–4/2 and Zh1–15 in the Zharmasaur Arc are euhedral or slightly round, and are commonly 30–150 μm long (Fig. 3c, d). CL images of these zircon grains display that they have mostly oscillatory zoning, which together with the evidence of moderate luminescence and Th/U ratios (0.31–0.95, Supplementary Table 1), indicates their magmatic origin. Sample Zh2–4/2 is characterized by a major zircon age population of 306–354 Ma, with a major peak at ~ 325 Ma (Fig. 4e, f), and $\varepsilon_{\text{Hf}}(t)$ values of these zircons are in a narrow range of 10.3 to 16.5 (Fig. 5b). Additional two zircon grains yield

relatively young ages of 291 Ma and 297 Ma. These Permian ages are younger than the Carboniferous deposition age of analyzed samples, and we consider that their U–Pb system may have been disturbed subsequently. Zircon grains from sample Zh1–15 give U–Pb ages from 332 Ma to 371 Ma, with a prominent peak at ~ 338 Ma (Fig. 4g, h).

Sample Ch–02–17 from the Char Zone contains stubby to prismatic zircons with low to high luminescence and relatively small sizes (40–100 μm long, Fig. 3e). They exhibit clear oscillatory zoning with relatively high Th/U ratios of 0.21–1.32 (Supplementary Table 1), indicating an igneous origin. 47 zircon grains from this sample give U–Pb ages in a narrow age range of 312–376 Ma, peaked at ~ 321 Ma and ~ 340 Ma (Fig. 4i, j). Additional one grain yields a Silurian age of 431 Ma.

4. Discussion

The closure of the Ob-Zaisan Ocean and the related arc amalgamation are imperative for the tectonic reconstruction of the western CAO. Our five samples for detrital zircon dating and Hf isotope analysis across the Char Shear Zone have significant implications for the location and time of the closure of the Ob-Zaisan Ocean.

4.1. Provenance of detrital zircons

Two Carboniferous samples (8–13–10/3 and K16–2) of the Irtysh-Zaisan Complex to the north of the Char Shear Zone show similar detrital zircon populations, and both are dominated by Devonian to Carboniferous detrital zircon ages (Fig. 4b, d). Combining detrital zircons from these two samples, total 124 analyses define four major age peaks of ~ 328 Ma, ~ 347 Ma, ~ 374 Ma and ~ 487 Ma (Fig. 6a), and have a broad $\varepsilon_{\text{Hf}}(t)$ range of -6.3 – 20.1 , similar as ages and $\varepsilon_{\text{Hf}}(t)$ values of detrital zircons from the Chinese segment of the Irtysh-Zaisan Complex (Figs. 6c and 7a; Li et al., 2015, 2017, 2019; Song et al., 2020b). Devonian to Carboniferous volcanic rocks widely occur within the Rudny Altai (Saraev et al., 2012; Lobanov et al., 2014; Kuibida et al., 2020), and ~ 338 – 318 Ma granitic rocks have been recognized along the Rudny Altai (Glorie et al., 2012; Kuibida et al., 2013; Kruk et al., 2014; Zhang et al., 2020). These igneous rocks likely represent major source of Devonian to Carboniferous detrital zircons in the Irtysh-Zaisan Complex. In addition, ~ 445 – 353 Ma plutonic rocks widely intrude into the Chinese segment of the Altai-Mongolian Terrane (Fig. 1b; Wang et al., 2006; Yuan et al., 2007; Sun et al., 2009; Wang et al., 2010; Cai et al., 2011b; Yang et al., 2011; Wang et al., 2011; Tong et al., 2012; Yu et al., 2017), and may also contribute to the provenance of Devonian to Early Carboniferous detrital zircons in the Irtysh-Zaisan Complex. Early Paleozoic detrital zircons with the peak age of ~ 487 Ma and Precambrian detrital zircons (554 Ma, 619 Ma, 675 Ma, 786 Ma, 835 Ma, 942 Ma, and 2028 Ma) in the Irtysh-Zaisan Complex, overlap in time with detrital zircon populations of early Paleozoic clastic rocks from the Altai-Mongolian Terrane (Fig. 6e; Long et al., 2007), indicating that these detrital zircons may recycle from the Altai-Mongolian Terrane.

Two Carboniferous samples (Zh2–4/2 and Zh1–15) from the Zharmasaur Arc to the south of the Char Shear Zone are dominated by Devonian to Carboniferous detrital zircon grains (Fig. 4f, h). Total 98 analyses from these two samples yield a major age peak at ~ 335 Ma (Fig. 6b), which is compatible with detrital zircon ages of the Zharmasaur Arc in the West Junggar of NW China (Fig. 6d; Li et al., 2017). Devonian to Carboniferous arc-related magmatism is widespread within the Zharmasaur Arc (Han et al., 2006; Zhou et al., 2008, 2015; Chen et al., 2010; Choulet et al., 2012a; Yang et al., 2014; Liu et al., 2018), and we consider that these magmatic rocks are the major source of Devonian to Carboniferous detrital zircons. In addition, $\varepsilon_{\text{Hf}}(t)$ of detrital zircons from sample Zh2–4/2 is characterized by positive values of 10–19 (Fig. 5b), which is consistent with the range of $\varepsilon_{\text{Hf}}(t)$ values from coeval detrital zircons in West Junggar (NW China; Fig. 7b; Li et al., 2017), suggesting juvenile origin of the coeval magmatism in the whole Zharmasaur Arc.

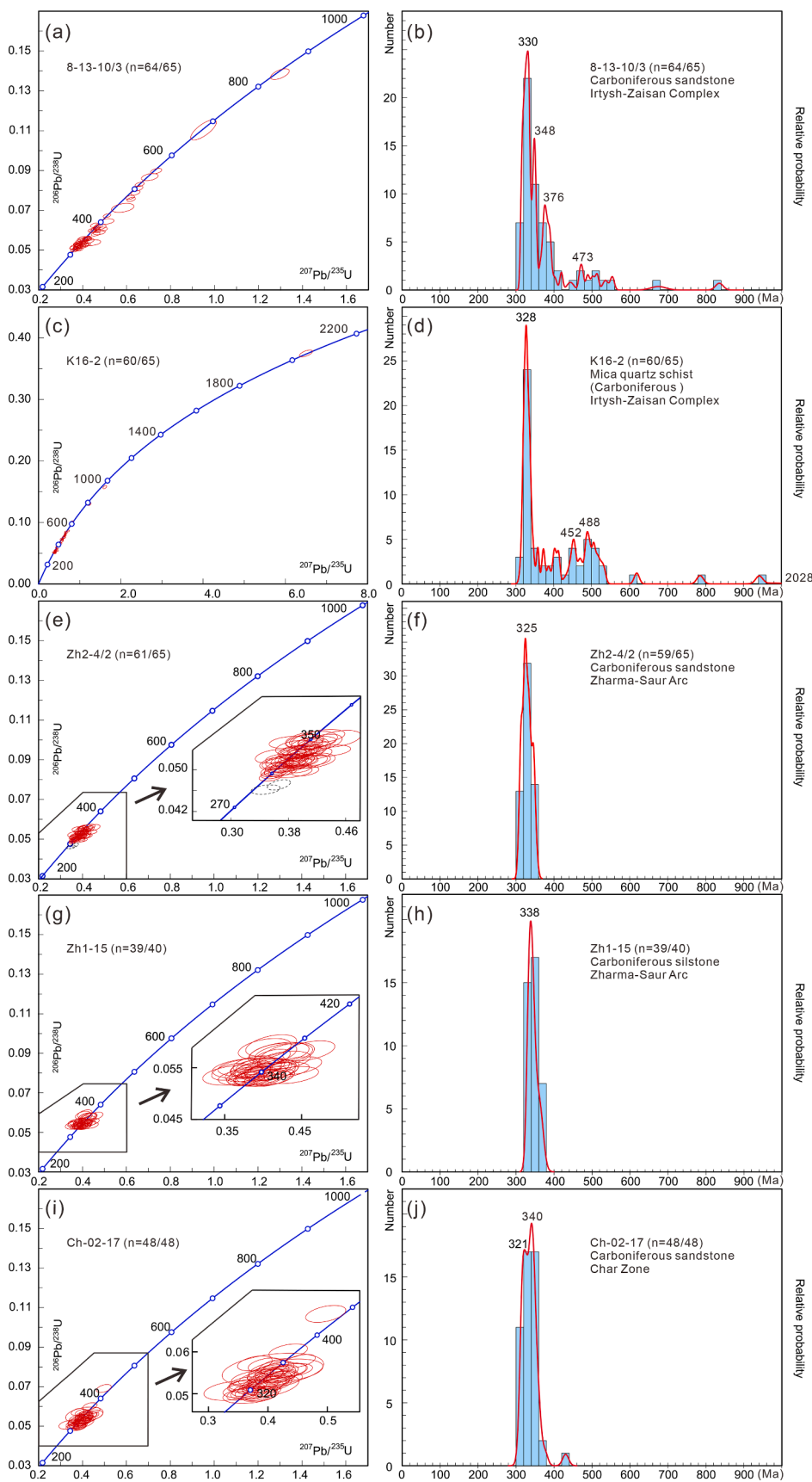


Fig. 4. U–Pb concordia plots (left) and age probability diagrams (right) of detrital zircons. The error is 1 σ level indicated by ellipses.

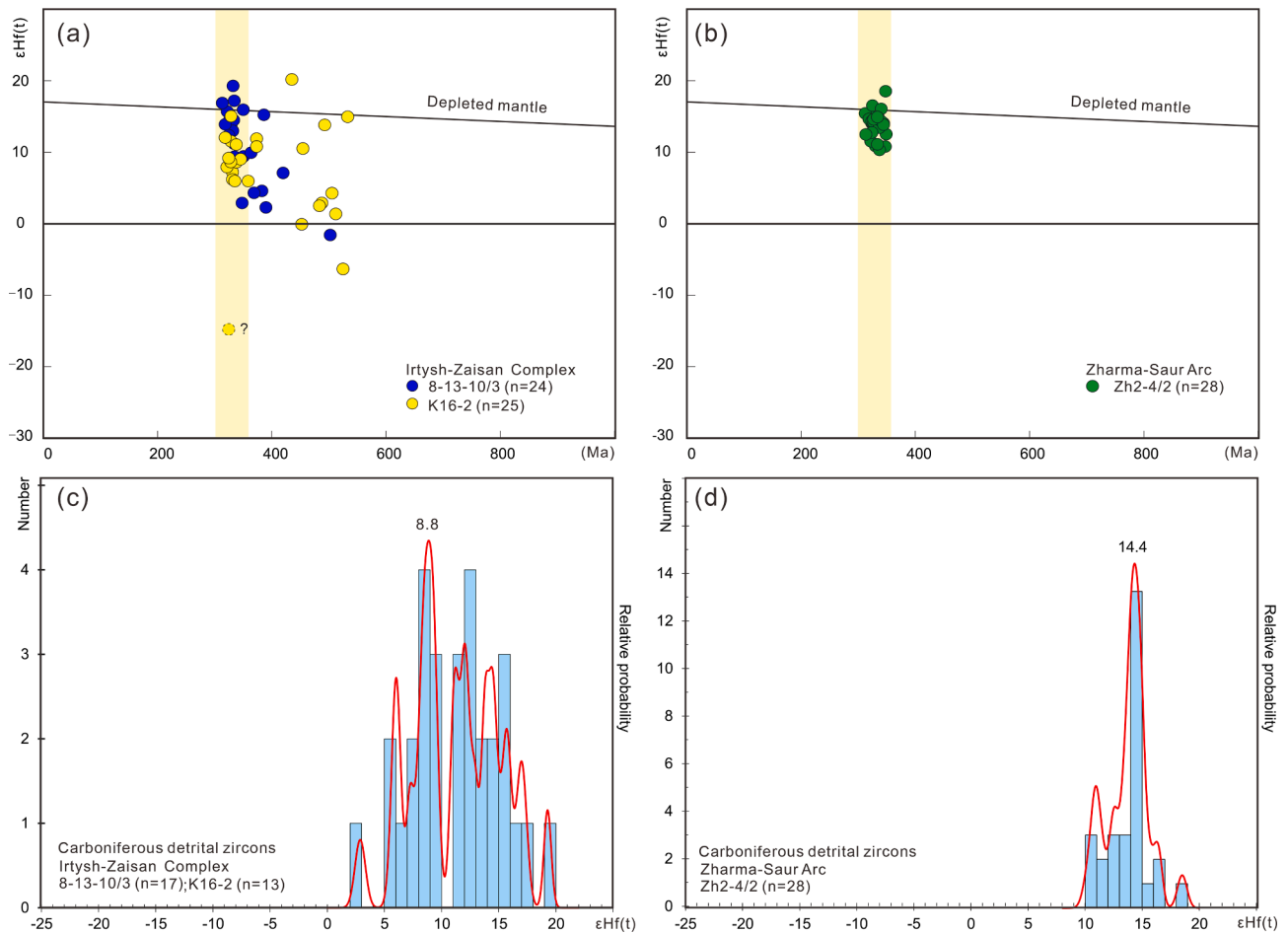


Fig. 5. (a-b) Diagram of $\epsilon_{\text{Hf}}(t)$ values versus crystallizing ages for detrital zircons in this study. (c-d) Probability diagrams of $\epsilon_{\text{Hf}}(t)$ values of Carboniferous detrital zircons.

An additional sample (Ch-02-17) was collected along the Char Zone. Available structural work demonstrates that a series of narrow sinistral faults characterize the Char Zone (Fig. 1b; Vladimirov et al., 2008), and therefore rocks from the Char Zone may have been transported horizontally in response to strike-slip movement. Our results show a major Devonian to Carboniferous detrital zircon population for sample Ch-02-17 (Fig. 4j), which is comparable to two samples from the Zharma-Saur Arc (Fig. 6b, f). A Silurian zircon grain (~431 Ma) from this sample overlaps within error with ~428 Ma andesite in the Zharma-Saur Arc (Zhang et al., 2015a), indicating that ~431 Ma detrital zircon grain may source from the Silurian magmatism within the Zharma-Saur Arc. Such data sets are consistent with the recent studies that show similar detrital zircon age populations of Carboniferous sedimentary rocks from the Char Zone as coeval clastic rocks in the Zharma-Saur Arc (Choulet et al., 2012a; Li et al., 2017; Safonova et al., 2021; Song et al., 2020a), which seems to indicate that clastic rocks within the Char Zone may tectonically belong to the Zharma-Saur Arc.

4.2. Tectonic implications

4.2.1. The tectonic boundary of the orogenic system around the Siberian margin and the Kazakhstan collage in eastern Kazakhstan

Our results show various provenances of sedimentary rocks across the Char Shear Zone. This supports the previous idea that the Char Shear Zone represents the tectonic boundary between the orogenic system around the Siberian margin and the Kazakhstan collage (Fig. 1b; Şengör et al., 1993; Travin et al., 2001; Buslov et al., 2004; Vladimirov et al., 2008; Glorie et al., 2012). The Carboniferous sedimentary rocks

(samples 8-13-10/3 and K16-2) from the Irtysh-Zaisan Complex (the Kazakhstan segment) to the northeast of the Char Shear Zone, contain abundant pre-Devonian detrital zircons peaked at ~488 Ma (Fig. 6a), similar to the detrital zircon populations from the Chinese segment of the Irtysh-Zaisan Complex (Fig. 6c; Li et al., 2015, 2017, 2019; Song et al., 2020b). On the contrary, the Carboniferous clastic rocks (sample Zh1-15 and Zh2-4/2) from the Zharma-Saur Arc to the southwest of the Char Shear Zone, are dominated by Devonian to Carboniferous detrital zircons that show a major age peak at ~335 Ma (Fig. 6b), which is consistent with the detrital zircon ages from the Zharma-Saur Arc in the West Junggar (NW China) (Fig. 6d; Choulet et al., 2012a; Li et al., 2017; Song et al., 2020a). In addition, the Carboniferous detrital zircons from the Zharma-Saur Arc have a narrow range of $\epsilon_{\text{Hf}}(t)$ values (around 10–19, Fig. 5c), which is distinct from the coeval detrital zircons from the Irtysh-Zaisan Complex that exhibit a relatively broad range (about 4–19, Fig. 5d), indicating different sources of the Carboniferous detrital zircons in the Zharma-Saur Arc and the Irtysh-Zaisan Complex.

On a larger scale, the Char Shear Zone extends into NW China, where it is merged with the Irtysh Shear Zone (Fig. 1b), with the latter representing the tectonic boundary of the Siberian margin (Chinese Altai) and the Kazakhstan collage (West Junggar) in NW China (Qu and Zhang, 1991, 1994; Li et al., 2017). Ophiolitic rocks have been recognized along both the Chinese segment of the Irtysh Shear Zone and the Char Shear Zone in eastern Kazakhstan (Buslov et al., 2001; Vladimirov et al., 2008; Volkova et al., 2008; Safonova et al., 2012, 2018; Wang et al., 2012; Song et al., 2020b), which together with ~444–429 Ma blueschists and eclogites within these shear zones (Buslov et al., 2001; Volkova et al., 2008) supports the interpretation with the Irtysh-Char Shear Zone as a

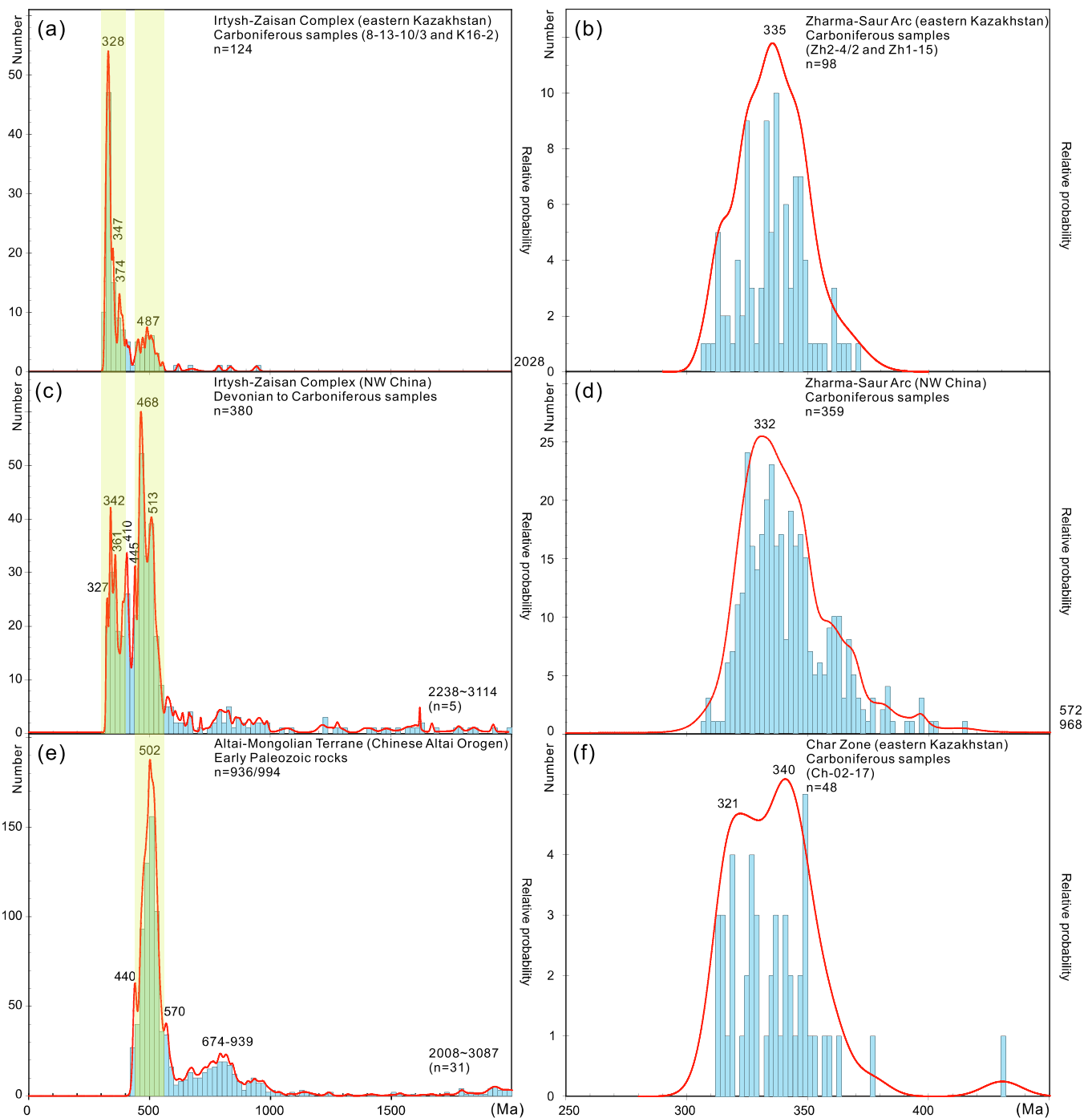


Fig. 6. (a-b) Age probability diagrams of all detrital zircon grains from the Irtysh-Zaisan Complex and the Zharma-Saur Arc in eastern Kazakhstan. (c-d) Age probability diagrams of detrital zircon grains from the Chinese segment of the Irtysh-Zaisan Complex (Li et al., 2015, 2017, 2019; Song et al., 2020b) and the Zharma-Saur Arc in the West Junggar of NW China (Choulet et al., 2012a; Li et al., 2017; Song et al., 2020a). (e) Age probability diagrams of detrital zircons from early Paleozoic rocks of the Chinese Altai Orogen (Li et al., 2019 and reference therein). (f) Age probability diagrams of detrital zircons from the Char Zone.

major suture zone of the Siberian margin and the Kazakhstan collage.

4.2.2. The collision of the Siberian margin and Kazakhstan collage in eastern Kazakhstan

The collisional front of the Siberian margin and the Kazakhstan collage is represented by the Irtysh-Zaisan Complex and the Zharma-Saur Arc/Char Zone (Fig. 1a). Buslov et al., (2004) considered that the Irtysh-Zaisan Complex in eastern Kazakhstan was probably deposited in a passive margin setting. However, the recent work along the Chinese segment of the Irtysh-Zaisan Complex suggests that it was probably

developed within a late Paleozoic accretionary wedge given the mixed occurrence of oceanic plate stratigraphy with continent-marginal clastic rocks (Briggs et al., 2007; Li et al., 2015, 2017; Xiao et al., 2015; Chen et al., 2019). Two Carboniferous samples (samples 8–13–10/3 and K16–2) from the Irtysh-Zaisan Complex in eastern Kazakhstan are characterized by abundant Carboniferous detrital zircons. These ages are close to the deposition time of the analyzed samples and are consistent with a convergent setting pattern on a cumulative probability plot (Fig. 8), thus supporting an accretionary wedge origin of the Irtysh-Zaisan Complex. The late Paleozoic arc-related magmatism was widely recognized within

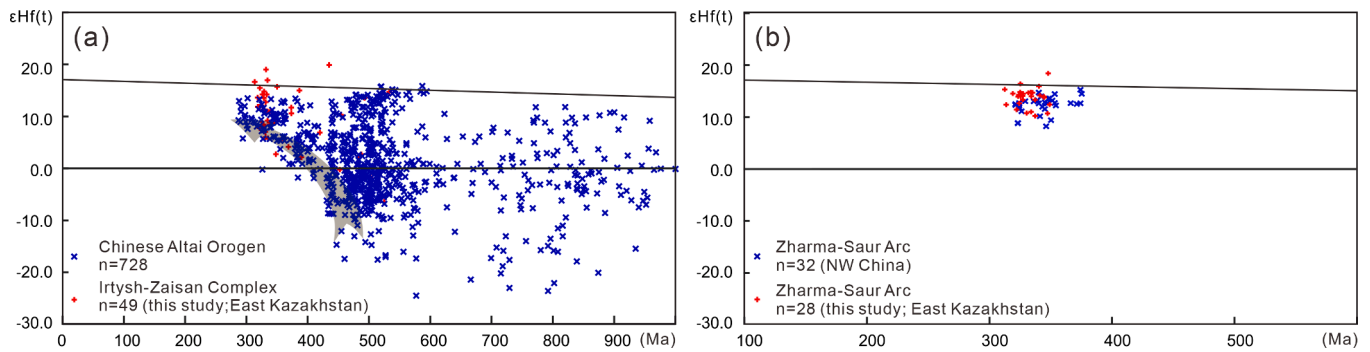


Fig. 7. (a) Diagram of $\epsilon_{\text{Hf}}(t)$ values versus crystallizing ages for detrital zircons from the Chinese Altai Orogen in NW China and the Irtysh-Zaisan Complex in eastern Kazakhstan (Li et al., 2019 and reference therein). The evolution of $\epsilon_{\text{Hf}}(t)$ towards positive values after ~ 400 Ma was interpreted to represent an episode of trench retreat (Li et al., 2019). (b) Diagram of $\epsilon_{\text{Hf}}(t)$ values versus crystallizing ages for detrital zircons from the Zharma-Saur Arc in eastern Kazakhstan and NW China (West Junggar) (Li et al., 2017).

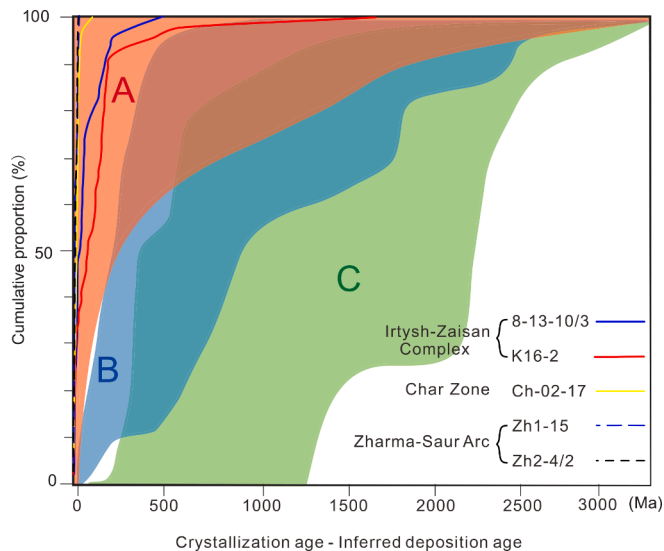


Fig. 8. Cumulative proportion curves for the difference of the crystallization age of detrital zircons with the depositional age of sedimentary rocks (Crystallization age-inferred deposition age). The plot of the general fields for convergent (A), collisional (B), and extensional (C) tectonic settings is based on Cawood et al. (2012).

the Rudny Aitai and Altai-Mongolian Terrane farther north (Vladimirov et al., 2008; Chai et al., 2009; Wan et al., 2010; Wang et al., 2011; Cai et al., 2012a; Kuibida et al., 2019; Zhang et al., 2020). The coeval development of magmatic arc and the Irtysh-Zaisan accretionary complex suggests a late Paleozoic arc-trench subduction system around the Siberian margin, associated with the northward subduction of the Ob-Zaisan oceanic plate. In addition, Carboniferous sedimentary rocks (samples Zh1-15 and Zh2-4/2) of the Zharma-Saur Arc are dominated by Devonian to Carboniferous detrital zircon grains (Fig. 4f, h) and exhibit a convergent setting pattern on a cumulative probability plot (Fig. 8), which is consistent with the previous interpretation that the Zharma-Saur Arc was developed in an intra-oceanic arc setting (Chen et al., 2017). Buslov et al. (2001) and Choulet et al. (2015) considered that the Zharma-Saur Arc was developed via southward subduction over the Bozshakol-Chingiz Arc (Fig. 1a). This interpretation is not preferred given the lack of Ordovician detrital zircons in the Zharma-Saur arc (Fig. 1b), which widely occurs within the Bozshakol-Chingiz Arc (Choulet et al., 2012a; Song et al., 2020a). Alternatively, the Zharma-Saur Arc may represent an independent late Paleozoic intra-oceanic arc (Safonova et al., 2017, 2021) separated from the Bozshakol-Chingiz Arc by an oceanic basin.

The collisional time of the Siberian margin and the Kazakhstan collage in eastern Kazakhstan can be constrained by our detrital zircon data. Our results demonstrate the absence of pre-Devonian detrital zircon grains within the Carboniferous strata of the Zharma-Saur Arc and the Char Zone, but such zircons widely exist within the Irtysh-Zaisan Complex (Section 4.1). This indicates that the Carboniferous strata of the Zharma-Saur Arc and the Char Zone did not receive detrital zircons from the orogenic system around the Siberian margin, and the Ob-Zaisan Ocean still bounded the Siberian margin with the Kazakhstan collage during the deposition of the Carboniferous strata of the Zharma-Saur Arc and the Char Zone. Three samples of Zh2-4/2, Zh1-15, and Ch-02-17 from the Zharma-Saur Arc and the Char Zone give the youngest age peaks of ~ 325 Ma, ~ 338 Ma, ~ 321 Ma, respectively. We consider that the youngest detrital zircon age peak of ~ 321 Ma from sample Ch-02-17 (Fig. 4j) within the Carboniferous strata of the Char Zone can give a maximum time constraint for the closure of the Ob-Zaisan Ocean. This timing constraint is compatible with the detrital zircon age data from NW China (West Junggar), which suggest that the Ob-Zaisan Ocean in NW China was closed after ~ 323 Ma (Li et al., 2017).

The youngest constraint for the closure of the Ob-Zaisan Ocean in eastern Kazakhstan can be determined by the activity time of the sinistral Char Shear Zone. This shear zone represents the tectonic boundary of the Siberian margin and the Kazakhstan collage in eastern Kazakhstan, and both systems must be collided prior to the sinistral shearing. The direct dating of the Char Shear Zone has not been undertaken till now, but this shear zone merges into the Irtysh Shear Zone in NW China. Therefore, both Char and Irtysh shear zones were supposed to be active simultaneously. The $^{40}\text{Ar}/^{39}\text{Ar}$ dating on *syn*-shearing minerals along the Irtysh Shear Zone constrains an age range of ~ 286 – 265 Ma for the sinistral strike-slip deformation (Fig. 1b; Travin et al., 2001; Laurent-Charvet et al., 2003; Buslov et al., 2004; Zhang et al., 2012; Li et al., 2017; Hu et al., 2020), which together with our detrital zircon data suggests that the closure of the Ob-Zaisan Ocean in eastern Kazakhstan occurred within a period of ~ 321 – 286 Ma, similarly as the closure of this oceanic basin in NW China (Li et al., 2015, 2017; Hu et al., 2020).

The development of a series of sinistral strike-slip shear zones (Char, Irtysh and North-East shear zones/faults) along the tectonic boundary of the Siberian margin and the Kazakhstan collage in the Permian suggests the oblique convergence of the both systems in this period (Fig. 9; Qu and Zhang, 1994; Buslov et al., 2003, 2004; Vladimirov et al., 2008; Kuibida et al., 2019). Li et al. (2015, 2017) and Hu et al. (2020) recognized early Permian sub-horizontal foliations and related \sim NW-SE orogen-parallel stretching lineations along the southern Chinese Altai Orogen (the southernmost segment of the *peri*-Siberian system in NW China), and suggested an episodic orogen-parallel extensional event prior to the oblique convergence, but post to the initial closure of the Ob-Zaisan Ocean. It remains enigmatic whether such earlier crustal

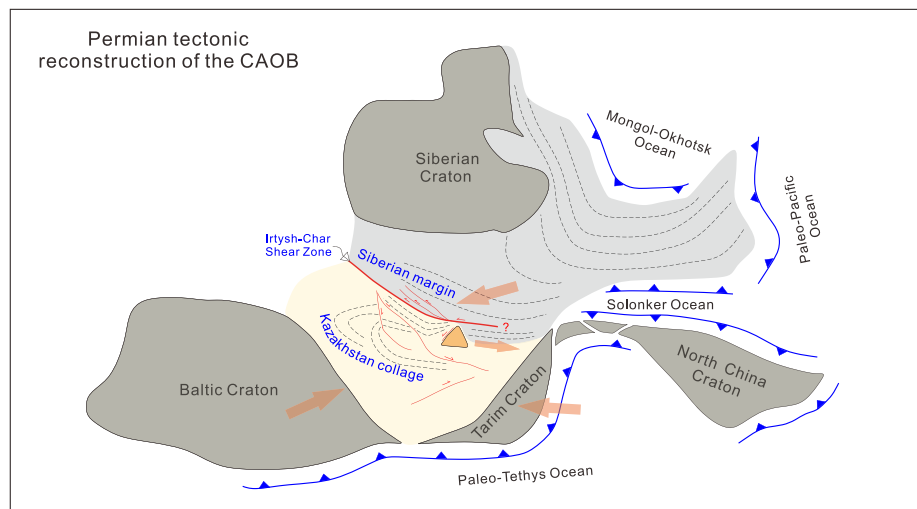


Fig. 9. Simplified tectonic reconstruction of the Central Asian Orogenic Belt in the Permian (after Xiao et al., 2015; Cunningham and Zhang, 2021). The latest Paleozoic strike-slip faulting was compatible with the eastward escape of the orogenic materials of the CAOB between the Siberian margin and the Tarim Craton (Li et al., 2015, 2017; Hu et al., 2020; He et al., 2021).

thickening and orogen-parallel extension also characterized early stage of the collision of the Siberian margin and Kazakhstan collage in eastern Kazakhstan, given the lack of detailed structural synthesis across the Irtysh-Zaisan Complex and the Zharma-Saur Arc in eastern Kazakhstan. In the scale of the whole western CAOB, the sinistral shearing along the Char, Irtysh, and North-East shear zones/faults occurred simultaneously with a series of dextral shear zones in the Tianshan Orogen farther south (Fig. 9; Shu et al., 1999; Laurent-Charvet et al., 2003; Lin et al., 2009; Wang et al., 2009; Cai et al., 2012b; Konopelko et al., 2013; Han and Zhao, 2018; Li et al., 2020, 2021; He et al., 2021). Such kinematic patterns suggest an eastward escape of orogenic segments between the Siberian margin and the Tianshan Orogen, which may link to the coeval convergence of the Siberian, Baltic, and Tarim cratons in the Permian (Li et al., 2015; Hu et al., 2020; He et al., 2021; Zhang et al., 2021).

5. Conclusions

Geochronological and Hf isotope analyses of detrital zircon samples across the Char Shear Zone provide crucial constraints for the collision of the orogenic system around the Siberian margin and the Kazakhstan collage, which are represented by the Irtysh-Zaisan Complex and the Zharma-Saur Arc/Char Zone, respectively. Carboniferous clastic rocks from the Irtysh-Zaisan Complex contain abundant early Paleozoic detrital zircons and minor Precambrian detrital zircons, which do not appear in Carboniferous sedimentary rocks of the Zharma-Saur Arc and the Char Zone. Devonian to Carboniferous detrital zircons exist in all sampling tectonic units across the Char Shear Zone, but predominant $\epsilon_{\text{Hf}}(t)$ values of the Carboniferous detrital zircons from the Zharma-Saur Arc show a narrower range (~10–17), which is different from coeval detrital zircons within the Irtysh-Zaisan Complex (~5–18). The distinct sedimentary provenances for the Irtysh-Zaisan Complex and the Zharma-Saur Arc/Char Zone confirm that their tectonic boundary lies along the Char Shear Zone in eastern Kazakhstan. Given that the Late Carboniferous sedimentary rocks from the Zharma-Saur and the Char Zone do not contain detritus originated from the Siberian marginal orogenic system, we conclude that the Ob-Zaisan Ocean between the Siberian margin and the Kazakhstan collage was closed after the Late Carboniferous (<321 Ma) deposition in the Zharma-Saur Arc and the Char Zone.

CRedit authorship contribution statement

Wanwan Hu: Conceptualization, Methodology, Validation,

Investigation, Writing – original draft, Writing – review & editing. **Pengfei Li:** Conceptualization, Methodology, Validation, Writing – review & editing, Supervision, Funding acquisition. **Min Sun:** Conceptualization, Validation, Writing – review & editing, Funding acquisition. **Inna Safonova:** Investigation, Validation, Writing – review & editing, Funding acquisition. **Yingde Jiang:** Writing – review & editing, Funding acquisition. **Chao Yuan:** Writing – review & editing, Funding acquisition. **Pavel Kotler:** Investigation, Funding acquisition.

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Acknowledgments

This study was funded by the National Natural Science Foundation of China (41872222 and 42021002), the international partnership program of the Chinese Academy of Sciences (132744KYSB20200001 and 132744KYSB20190039), Hong Kong Research Grant Council (HKU17302317), the Guangdong Province (project 2019QN01H101), and the Russian Science Foundation (project 21-77-20022 for Inna Safonova, tectonics; project 20-7710051 for Pavel Kotler, regional geology). We acknowledge Dan Wu and Le Zhang for the assistance of the analytical work. The manuscript benefits from the comments by Dongfang Song and an anonymous reviewer. This is a contribution of Guangzhou Institute of Geochemistry, Chinese Academy of Sciences (GIG-CAS; No. IS-3091), Sobolev Institute of Geology and Mineralogy SB RAS (State Assignment # 033020190003), the Chemical Geodynamics Joint Laboratory between Hong Kong University and GIG-CAS as well as IGCP 662.

Appendix A. Supplementary material

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.jseae.2021.104978>.

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